

Evolution of the Deep and Bottom Waters of the Scotia Sea, Southern Ocean, during 1995–2005*

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ABSTRACT

The Southern Ocean hosts the formation of the densest layers of the oceanic overturning circulation and provides a climatically sensitive element of deep ocean ventilation. An oceanographic section across the eastern Scotia Sea occupied in 1995, 1999, and 2005 reveals significant variability in the deep and bottom waters of Southern Ocean origin. Warming ($\sim 0.1^{\circ}\text{C}$) of the warm midlayer waters in the Scotia Sea between 1995 and 1999 reversed through to 2005, reflecting changes seen earlier upstream in the Weddell Sea. The volume of deep waters with potential temperature less than 0°C decreased during 1995–2005, though such a reduction was only clear between 1995 and 1999 at the southern end of the section. The abyssal waters of the eastern Scotia Sea changed circulation between 1995 and 1999, with the dominant point of their entry to the basin shifting from the south to the northeast; by 2005, the former route had regained dominance. These changes are best explained by interannual variations in the deep waters exiting the Weddell Sea, superimposed on a longer-term (decadal) warming trend. The interannual variations are related to changes in the strength of the Weddell Gyre, reflecting large-scale atmospheric variability that may include the El Niño–Southern Oscillation phenomenon. The Scotia Sea is the most direct pathway for dense waters of the overturning circulation emanating from the Weddell Sea to fill much of the World Ocean abyss. The regional changes reported here have the potential to affect the climatically significant ventilation of the global ocean abyss.

1. Introduction

Immediately east of Drake Passage, the Scotia Sea is a basin with rough bottom topography and is

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separated from the Weddell Sea to the south by the South Scotia Ridge (Fig. 1). The horizontal zonation of water masses in the Scotia Sea is dominated by the Antarctic Circumpolar Current (ACC; Orsi et al. 1995). This current has a banded structure, with the main fronts being (from north to south) the Sub-Antarctic Front, Polar Front, and Southern ACC Front. The southern limit of the ACC is marked by the Southern Boundary, south of which are found the waters of the Weddell–Scotia Confluence. These are characterized by low vertical gradients at middepths resulting from the injection of dense waters that spill off the shelf close to the tip of the Antarctic Peninsula (Whitworth et al. 1994).

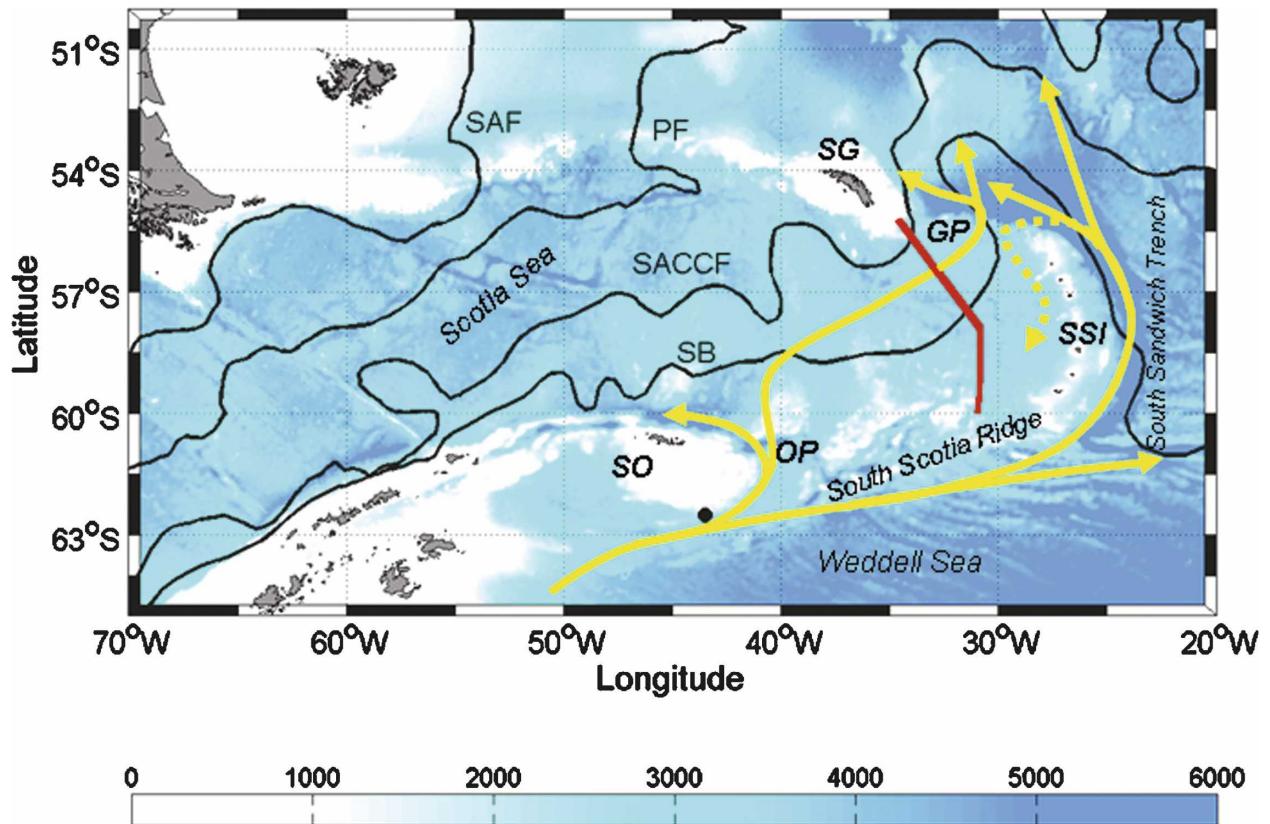


FIG. 1. Location of repeat section used here (red line) and position of the CORC ARCHES M2 mooring (black dot) immediately southeast of the South Orkneys (SO), with bathymetry (Smith and Sandwell 1997) color shaded (m). Antarctic Circumpolar Current fronts marked schematically: Subantarctic Front (SAF), Polar Front (PF), and Southern ACC Front (SACCF) positions from Boehme et al. (2008, manuscript submitted to *J. Geophys. Res.*), and Southern Boundary (SB) from Orsi et al. (1995). Marked topographic features include South Georgia (SG), SO, Orkney Passage (OP), Georgia Passage (GP), and South Sandwich Islands (SSI). The primary route of spreading of Weddell Sea Deep Water through and around the Scotia Sea is shown schematically (yellow solid arrows), with the intermittent inflow of its densest component through GP also marked (yellow dashed arrow).

The Scotia Sea receives waters from both Drake Passage, via the ACC, and the Weddell Sea, via the northern limb of the Weddell Gyre. Water masses of Weddell Sea origin include warm deep water (WDW; Table 1), characterized by a potential temperature (θ) maximum near 500 m and a salinity (S) maximum near 800 m. WDW is derived from the Circumpolar Deep Water (CDW) of the ACC that is entrained into the eastern limb of the Weddell Gyre (Orsi et al. 1993), and thereafter advected cyclonically within the gyre.

Waters of Weddell Sea origin also include Weddell Sea Deep Water (WSDW; Table 1), the densest water mass found in the Scotia Sea. WSDW is an essential component of the Antarctic Bottom Water (AABW) that forms the abyssal layer of much of the world's oceans. WSDW is formed by the mixing of dense shelf waters with variants of WDW on the shelves and slopes of the western and southwestern Weddell Sea (Foster and Carmack 1976; Gill 1973; Gordon 1998). Different

types of WSDW are found in the Weddell Sea reflecting distinct formation processes and regions. The outer rim of the gyre features relatively fresh and oxygen-rich WSDW compared with that found closer to the center

TABLE 1. Scotia Sea and Weddell Sea water masses of Southern Ocean origin discussed in this paper; γ'' is the neutral density variable of Jackett and McDougall (1997; kg m^{-3}).

Water mass	Abbreviation	Identifiers
Warm deep water	WDW	θ_{\max} near 500 m S_{\max} near 800 m $\gamma'' < 28.20$ in the ACC
Weddell Sea Deep Water	WSDW	$28.26 < \gamma'' < 28.40$
Upper Weddell Sea Deep Water	Upper WSDW	$28.26 < \gamma'' < 28.31$
Lower Weddell Sea Deep Water	Lower WSDW	$28.31 < \gamma'' < 28.40$
Weddell Sea Bottom Water	WSBW	$\gamma'' > 28.40$

of the gyre. The enhanced ventilation is due to injection of relatively fresh shelf water found near the northern end of the Antarctic Peninsula (Gordon et al. 2001). WSDW is underlain in the Weddell Sea by the denser Weddell Sea Bottom Water (WSBW; Table 1), which forms via similar processes to WSDW but contains a higher proportion of dense shelf waters. WSBW is too dense to escape the Weddell Sea directly and must first mix upward into the WSDW density class. Within the Weddell Sea, WSDW can also incorporate recently ventilated deep waters imported from the Indian Ocean sector to the east (Meredith et al. 2000).

The Scotia Sea (Fig. 1) is the most direct route for WSDW to flow northward and contribute to the global overturning circulation (Locarnini et al. 1993). WSDW can enter the Scotia Sea via deep gaps in the South Scotia Ridge, of which the ~ 3200 -m-deep Orkney Passage is the deepest (Gordon et al. 2001; Naveira Garabato et al. 2002b), or via gaps in the South Sandwich Island arc (Meredith et al. 2001; Naveira Garabato et al. 2002a). The northward flow of WSDW on the western edge of the Orkney Passage derives from the more ventilated form of WSDW (Gordon et al. 2001). Upon entering the Scotia Sea, WSDW can spread westward toward the Drake Passage; however, there is little or no net flow of this water mass through the narrowest part of the passage into the Pacific Ocean (Gordon 1966; Sievers and Nowlin 1984; Nowlin and Zenk 1988). WSDW exits the Scotia Sea primarily in the vicinity of the ~ 3200 -m-deep Georgia Passage, southeast of South Georgia (Fig. 1).

Upon leaving the Scotia Sea, WSDW spreads northward into the Georgia Basin, whence it can continue northward into the Argentine Basin or recirculate southeastward back toward the Weddell Sea. Conventionally, WSDW is divided into upper and lower components in this region (Table 1; Arhan et al. 1999; Meredith et al. 2001). The division between Upper and Lower WSDW [established by Arhan et al. (1999), and used here] is such that Upper WSDW is able to flow through the Scotia Sea to reach the Georgia Basin, whereas Lower WSDW is topographically constrained to enter the Georgia Basin via the eastern side of the South Sandwich Islands (Fig. 1). The direction of abyssal Lower WSDW flow at Georgia Passage was inferred to have reversed between 1995 and 1999, with the flow being northeastward out of the Scotia Sea prior to 1995, but then changing to be southwestward into the Scotia Sea (Meredith et al. 2001).

Various changes have been reported in the properties of water masses originating from the Weddell Sea. Analysis of a long compilation of hydrographic data

revealed that the coldest WDW was associated with the Weddell Polynya years of the 1970s (Robertson et al. 2002). Subsequently, this water mass warmed at a rate of around $0.012 (\pm 0.007) ^\circ\text{C yr}^{-1}$ up to the 1990s. This warming was not compensated by a salinity increase, with the WDW becoming less dense as a result. The Scotia Sea exhibited a warming of WDW of around $0.1^\circ\text{--}0.2^\circ\text{C}$ between 1995 and 1999, as changes originating in the Weddell Sea arrived (Meredith et al. 2001); however, more recent data from close to the Greenwich meridian in the Weddell Sea revealed a cooling of around 0.05°C between 1998 and 2003 after the end of the warming signal (Fahrback et al. 2004). Tentative explanations for these changes include a change in CDW properties in the ACC (e.g., Gille 2002) combined with variations in the rate of inflow of this water to the Weddell Gyre.

Weddell Sea temperature profiles suggest warmer WSDW between 1500 and 3500 m in the 1990s compared with the 1970s (Robertson et al. 2002). However, high levels of variability in the data prevented identification of a well-defined trend. Relatively little change in WSDW properties in the Weddell Sea was evident during 1990–2003, with variations being close to the detection level (Fahrback et al. 2004). The areas from which data were collected for that study are, however, likely to emphasize basin-wide bulk changes in properties rather than changes in the recently ventilated forms of WSDW in the boundary currents at the periphery of the sea. The comparative stability of bulk WSDW properties contrasts with the continuous warming (by $\sim 0.025^\circ\text{C}$) experienced by WSBW between 1989 and 2003 (Fahrback et al. 2004).

Temperature changes in the bottom waters spreading northward in the Atlantic Ocean from the Weddell Sea are apparent all the way to the equator. Directly downstream from the Scotia Sea, in the Argentine Basin, a comparison of hydrographic data collected in 1989 to data from 1980–87 showed that the coldest abyssal layer warmed significantly in that interval (Coles et al. 1996). Also in the Argentine Basin, a comparison of 2005 and 1989 hydrographic section data revealed a further $\sim 0.05^\circ\text{C}$ warming of the abyssal layer (Johnson and Doney 2006). In the Vema Channel, the deepest conduit between the Argentine Basin and the Brazil Basin to the north, the coldest bottom waters have warmed by as much as $2.8 \times 10^{-3} ^\circ\text{C yr}^{-1}$ from 1991 through to 2006, after being relatively stable from 1972 to 1991 (Zenk and Morozov 2007). In the Brazil Basin, the abyssal layers warmed by $\sim 0.04^\circ\text{C}$ between 1989 and 2005 (Johnson and Doney 2006). Finally, at the equator, bottom θ values increased by nearly 0.1°C between 1993 and 1999 (Andri  et al. 2003). These changes re-

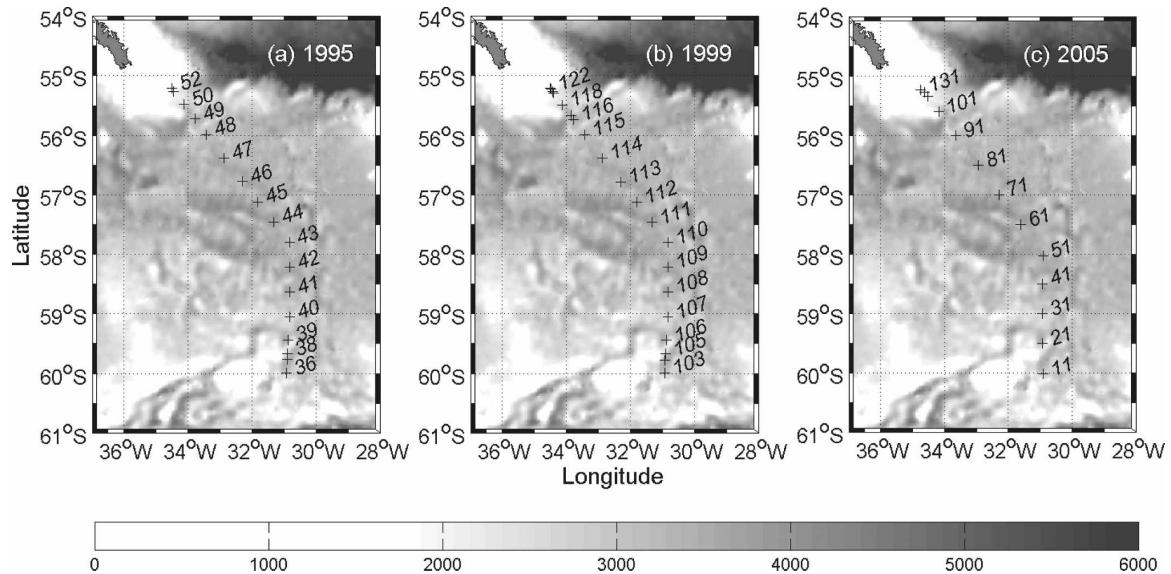


FIG. 2. Locations of CTD stations used here (see Fig. 1 for broader context). Data were collected during (a) 1995, (b) 1999, and (c) 2005. Background shading is depth (m).

late to abyssal temperatures, and while it is important to note that there are differences in bottom temperatures and lateral bottom temperature gradients between each of these locations, the order of magnitude and timing of the changes are suggestive of a decadal warming trend propagating northward in the South Atlantic.

In this paper, we use repeat hydrographic data to examine the evolution of the water mass properties and circulation of Weddell Sea-derived waters in the eastern Scotia Sea over a recent decade (1995–2005), and investigate the causes of the observed changes. Given these waters' fundamental contribution to the deepest layers of the global overturning circulation, it is important to monitor and ultimately identify the causes of their variability.

2. Methods

Hydrographic data were collected from three cruises in the eastern Scotia Sea that occupied virtually identical tracks (Fig. 2). Each of these ran northward from the South Scotia Ridge along $\sim 31^\circ\text{W}$, and then diverted at $\sim 58^\circ\text{S}$ toward the island of South Georgia (Fig. 1). The first section was conducted in April 1995 from the British Antarctic Survey vessel Royal Research Ship (RRS) *James Clark Ross*, and was a component of the World Ocean Circulation Experiment (WOCE) A23 section (Heywood and King 2002). The second section was conducted from the same vessel in April 1999 (Naveira Garabato et al. 2002a), and the

third was conducted in January 2005 from the National Oceanic and Atmospheric Administration (NOAA) ship *Ronald H. Brown*, as part of the U.S. Climate Variability and Predictability (CLIVAR)/CO₂ Repeat Hydrography Program (Johnson and Doney 2006).

Meredith et al. (2001) described in detail the methods and data for the first two of these sections. To summarize, the 1995 data were collected with a Neil Brown MKIIIb conductivity–temperature–depth (CTD) profiler, with S data being calibrated relative to the International Association for the Physical Sciences of the Oceans (IAPSO) P-series standard seawater using bottle salts collected with a 24-bottle rosette. The accuracy of this data is 0.001°C for θ and 0.002 for S . The 1999 data were collected and calibrated in the same manner, but using a Neil Brown MKIIIc CTD, with comparable accuracies achieved. The 2005 data were collected using a SeaBird Electronics (SBE) *9plus* CTD fitted with SBE3 temperature sensors and SBE4 conductivity sensors, and with discrete samples obtained using a 36-position rosette. The accuracy of the θ and S data from 2005 are 0.001°C and 0.002 or better, respectively (Johnson and Doney 2006).

3. Zonation and water masses

Vertical sections of θ , S , and neutral density (γ^n) are presented in Figs. 3–5 respectively, for each of the years. The major feature of horizontal zonation along this line is the Southern Boundary, found at $\sim 58.5^\circ\text{S}$ in the 1995 section, $\sim 59^\circ\text{S}$ in the 1999 section, and be-

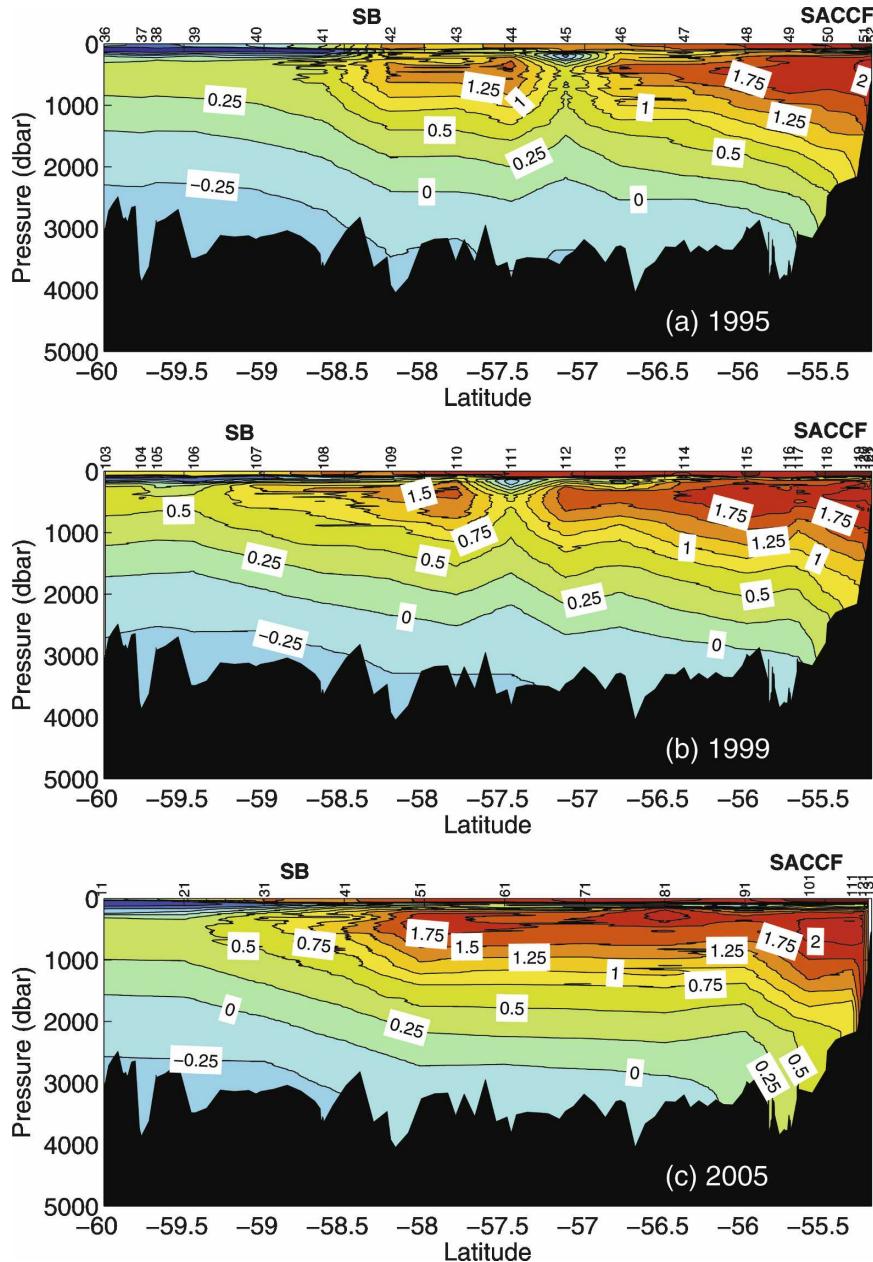


FIG. 3. Vertical meridional sections of potential temperature ($^{\circ}\text{C}$) from hydrographic sections (Fig. 1 shows location) occupied in (a) 1995, (b) 1999, and (c) 2005. Individual station locations marked along the upper horizontal axis (see also Fig. 2). Locations of the SB and SACCF are also marked.

tween 58.5° and 59°S in the 2005 section. Narrow eddy-like features containing water from south of the Southern Boundary are seen in the 1995 and 1999 sections at stations 45 and 111 (around 57° and 57.5°S , respectively; Meredith et al. 2001). The Southern ACC Front is present at the northern end of the sections, close to stations 50, 118, and 101 in 1995, 1999, and 2005, respectively.

North of the Southern Boundary, the sections are occupied by the ACC's poleward flank, in which vertical θ and S maxima indicate the cores of the upper and lower components of the CDW (Figs. 3 and 4, respectively). South of the Southern Boundary lies the Weddell–Scotia Confluence; this features WDW centered around 500 m, identifiable as the vertical θ maximum (Fig. 3). WDW of $\gamma'' < 28.20 \text{ kg m}^{-3}$ is constrained to

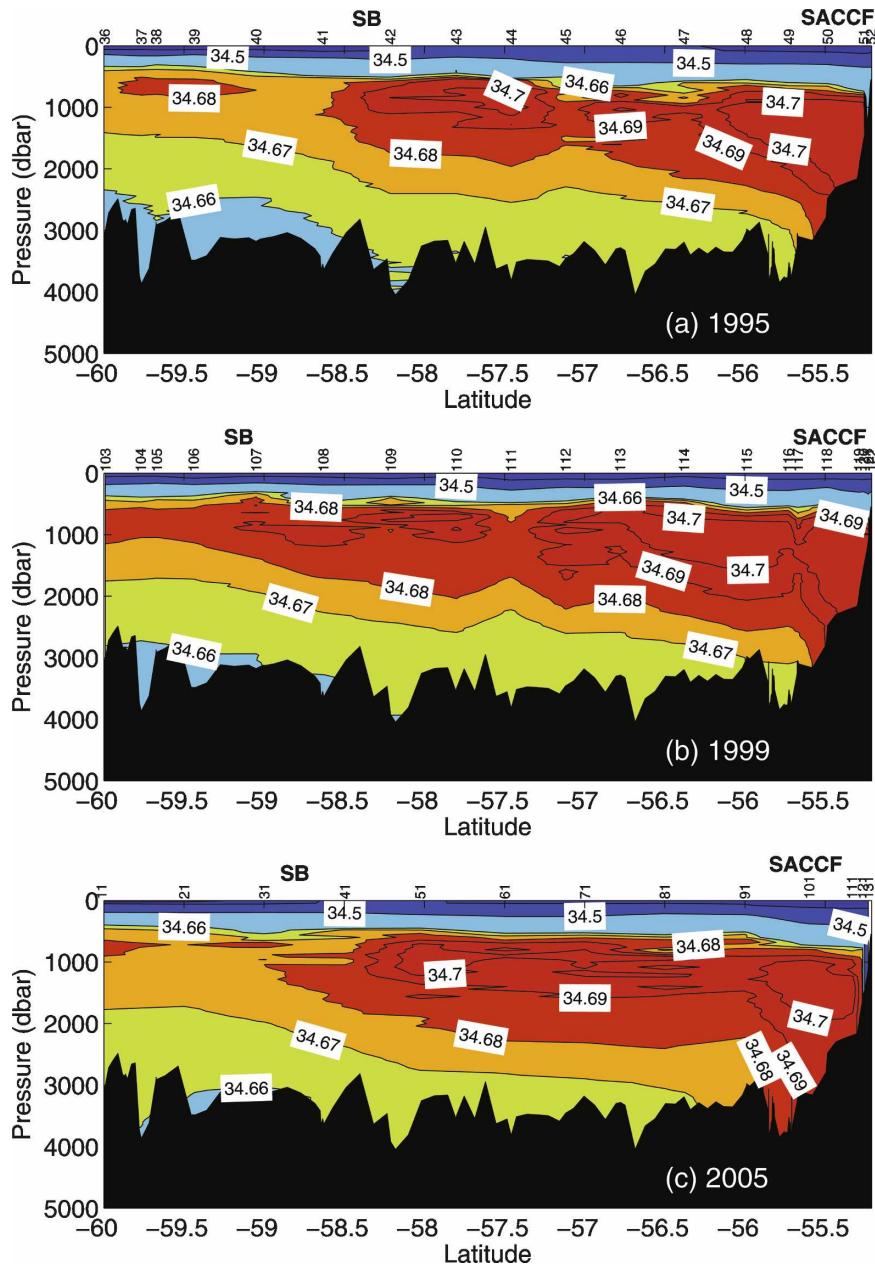


FIG. 4. Vertical meridional salinity (PSS-78) sections. Details follow Fig. 3. A nonuniform contour interval is used.

lie south of the Southern Boundary, except for its presence in the narrow eddy-like features noted above. A denser form of WDW is observed farther north at greater depths (>500 m), and has its origins in the modification of southeast Pacific Deep Water by mixing with Weddell Deep Water (Naveira Garabato et al. 2002a). Beneath the WDW and CDW on the section lies WSDW, here defined as waters of $\gamma'' > 28.26$ kg m^{-3} (Fig. 5). This layer resides below ~ 1800 m at the

southern end of the section, thins northward, and disappears at around 55.5° – 56° S.

4. Variability in water mass properties and circulation

a. Warm deep water variability

The WDW layer south of the Southern Boundary exhibits substantial variability in θ and S measured dur-

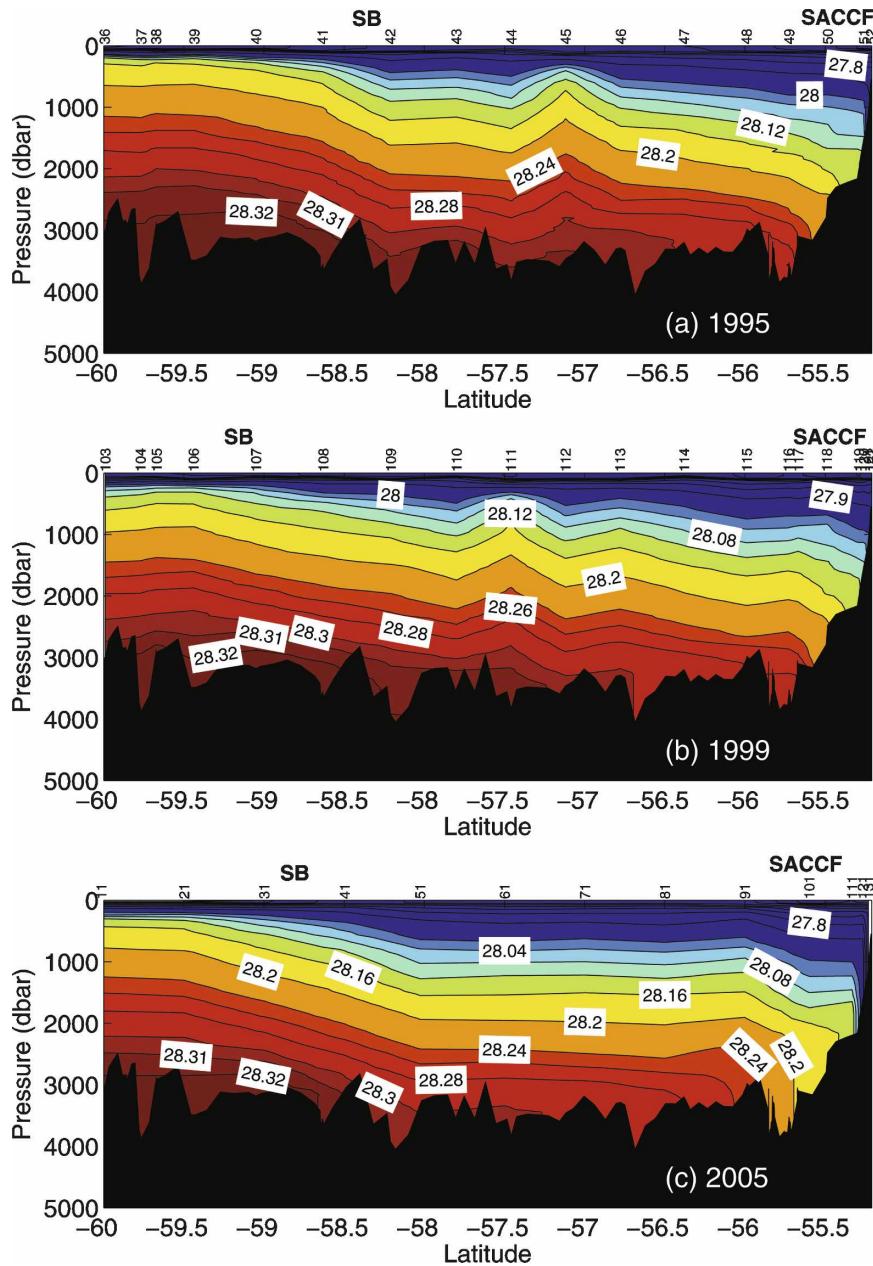


FIG. 5. Vertical meridional neutral density (γ^n ; kg m^{-3} ; Jackett and McDougall 1997) sections. Details follow Fig. 3. Note the change in contour interval for $\gamma^n > 28.28 \text{ kg m}^{-3}$.

ing the three cruises (Fig. 6). In 1995, the peak WDW θ value encountered was close to 0.41°C , with the majority of stations showing WDW core θ values of $0.35^\circ\text{--}0.40^\circ\text{C}$. In the 1999 occupation of the section, the WDW core θ values were all much higher than this, from 0.50° to 0.65°C . In the 2005 section, the WDW core θ had decreased again, with values below 0.45°C . These changes are reflected in WDW maximum S values also, which increased by ~ 0.004 between 1995 and 1999, before decreasing by a similar amount through to

2005 (Fig. 6). Note that there are some changes in the lateral position of the Southern Boundary in the sections (e.g., Fig. 3); however, our comparison considers only the waters south of this feature for each of the years, thus the changes we see cannot be due to such movements.

b. Weddell Sea Deep Water

Vertical sections of γ^n (Fig. 5) indicate that the volume of WSDW ($\gamma^n > 28.26 \text{ kg m}^{-3}$) decreased from

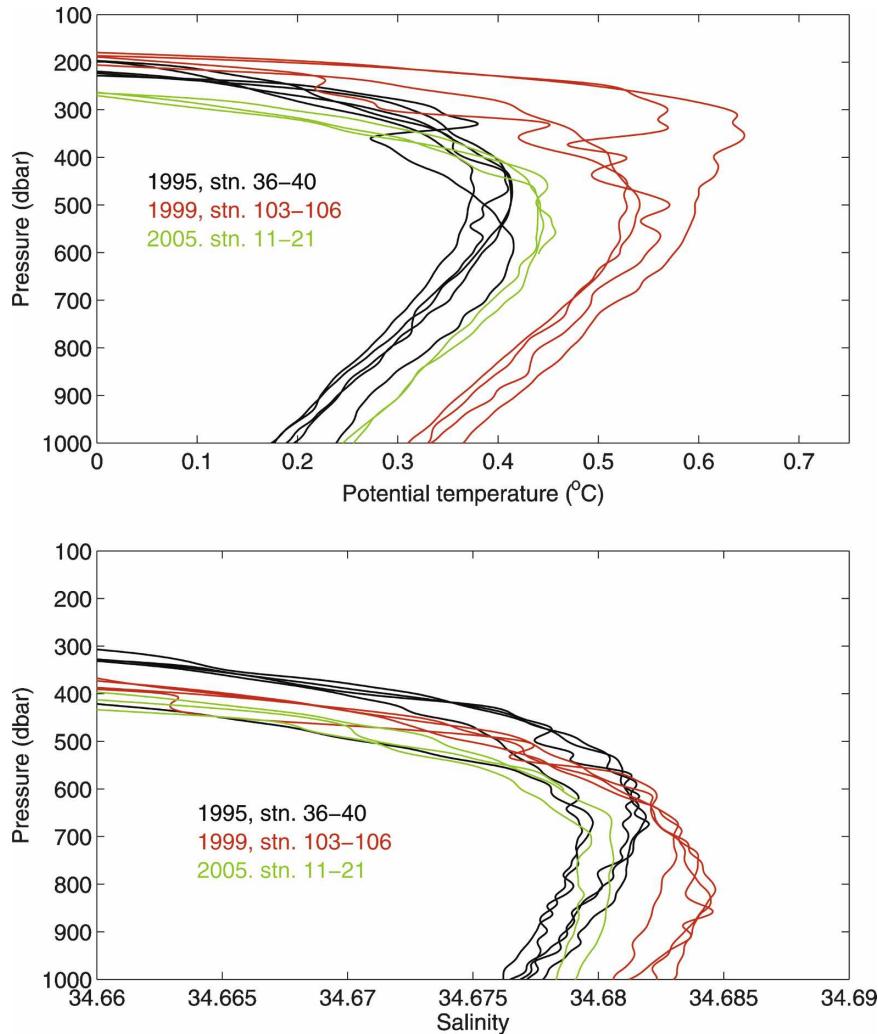


FIG. 6. (top) Potential temperature ($^{\circ}\text{C}$) and (bottom) salinity vs pressure south of the Southern Boundary for 1995 (black), 1999 (red), and 2005 (green). Profiles have been smoothed with a 20-dbar Butterworth filter. Station locations are shown in Fig. 2.

1995 to 1999 and decreased again from 1999 to 2005. In particular, surfaces of $\gamma^{\prime\prime} > 28.26 \text{ kg m}^{-3}$ appear to sink progressively, and their intersections with the seabed move southward, between 1995 and 2005. As an example, $\gamma^{\prime\prime} = 28.30 \text{ kg m}^{-3}$ intersected the seabed close to station 47 in 1995, that is, around 56.4°S . In 1999, this surface intersected the seabed around 56.8°S (station 113), and in 2005 its intersection was close to 57.4°S (station 61).

Although different station positioning among the years (most notably 2005; Fig. 2) makes station-by-station comparisons problematic, further indications of this shift can be obtained by considering vertical θ and S profiles from north of the Southern Boundary (Fig. 7). The deeper profile depths in 2005 are a reflection of the different station positioning in this year. This com-

plication notwithstanding, the 2005 stations are all warmer and saltier than the 1995 stations by around 0.1°C and 0.005 , respectively. The 1999 stations are intermediate, and overlap with both 1995 and 2005. The changes are largely density compensating, such that the θ - S curves from these stations are not significantly different within the accuracy level of the S measurements, except that the densest water present becomes progressively warmer and saltier (not shown).

South of the Southern Boundary there is a somewhat different pattern of change in θ within the WSDW (Fig. 8). The warming of around 0.05°C is present between 1995 and 1999, as discussed previously (Meredith et al. 2001). However, subsequent to this there are no further significant changes between 1999 and 2005, with θ values remaining constant on pressure surfaces. The

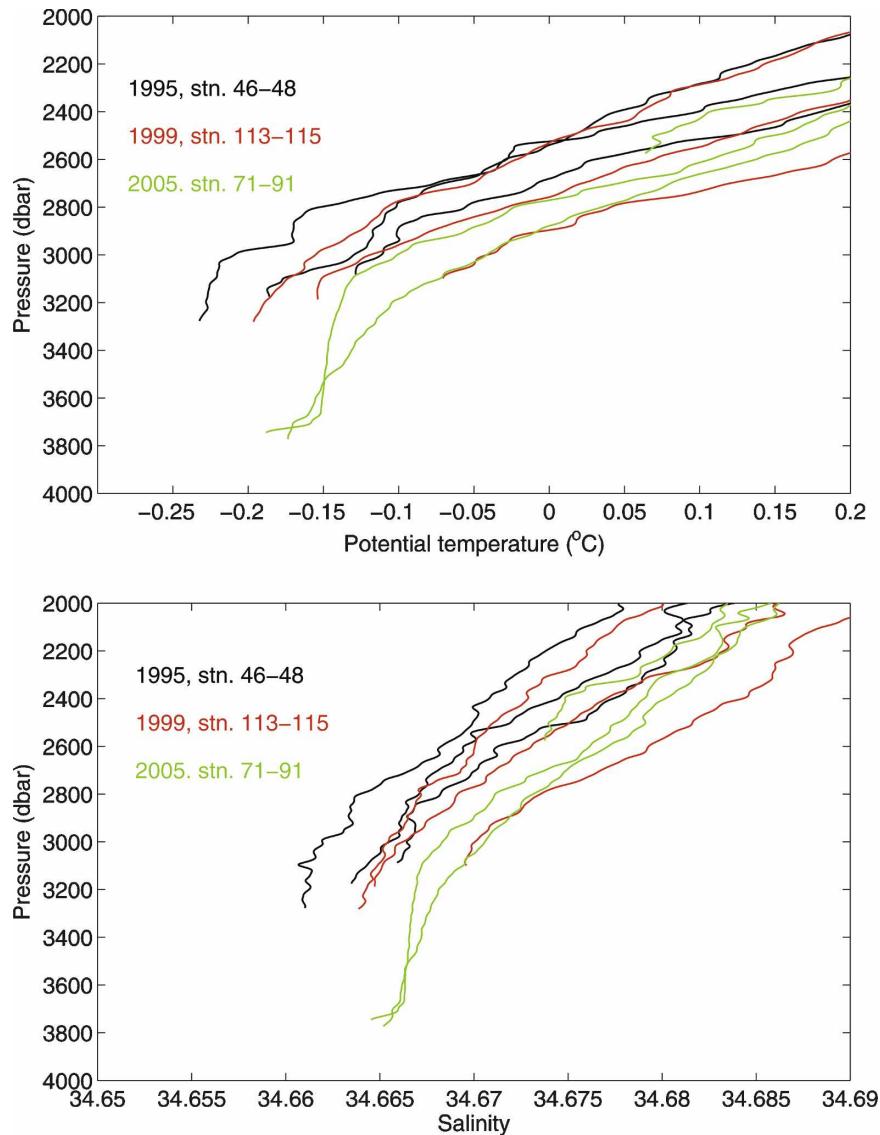


FIG. 7. (top) Potential temperature ($^{\circ}\text{C}$) and (bottom) salinity vs pressure for stations in the latitude range 56° – 57°S , north of the Southern Boundary and the narrow eddy feature seen in 1995 and 1999, for 1995 (black), 1999 (red), and 2005 (green). Profiles have been smoothed with a 20-dbar Butterworth filter. Station locations are shown in Fig. 2.

changes in S suggest a possible shift toward saltier WSDW during the sequence of observations (Fig. 8, lower panel), but it is small compared with the accuracy of the measurements.

c. Lower Weddell Sea Deep Water variability

As with WSDW as a whole there is a reduction in Lower WSDW volume during the sequence of measurements, though it is most notable between 1999 and 2005 for this layer. In 1995, the northern terminus of this water mass (defined as $\gamma^{\sigma} > 28.31 \text{ kg m}^{-3}$; Fig. 5)

was near 57°S , before moving slightly southward to $\sim 57.25^{\circ}\text{S}$ in 1999, then farther southward than 58°S in 2005. Within the Lower WSDW, an inflection in the θ – S relation at $\gamma^{\sigma} = 28.31 \text{ kg m}^{-3}$ is present in the 1999 data, with the curves below this level becoming saltier, but is absent from the 1995 and 2005 data (Fig. 9). The 1999 curves thus reflect the presence of an additional water mass in the Scotia Sea that was absent in the other years, namely, a thin, abyssal layer of comparatively salty Lower WSDW that entered the Scotia Sea via Georgia Passage and possibly other deep gaps in the

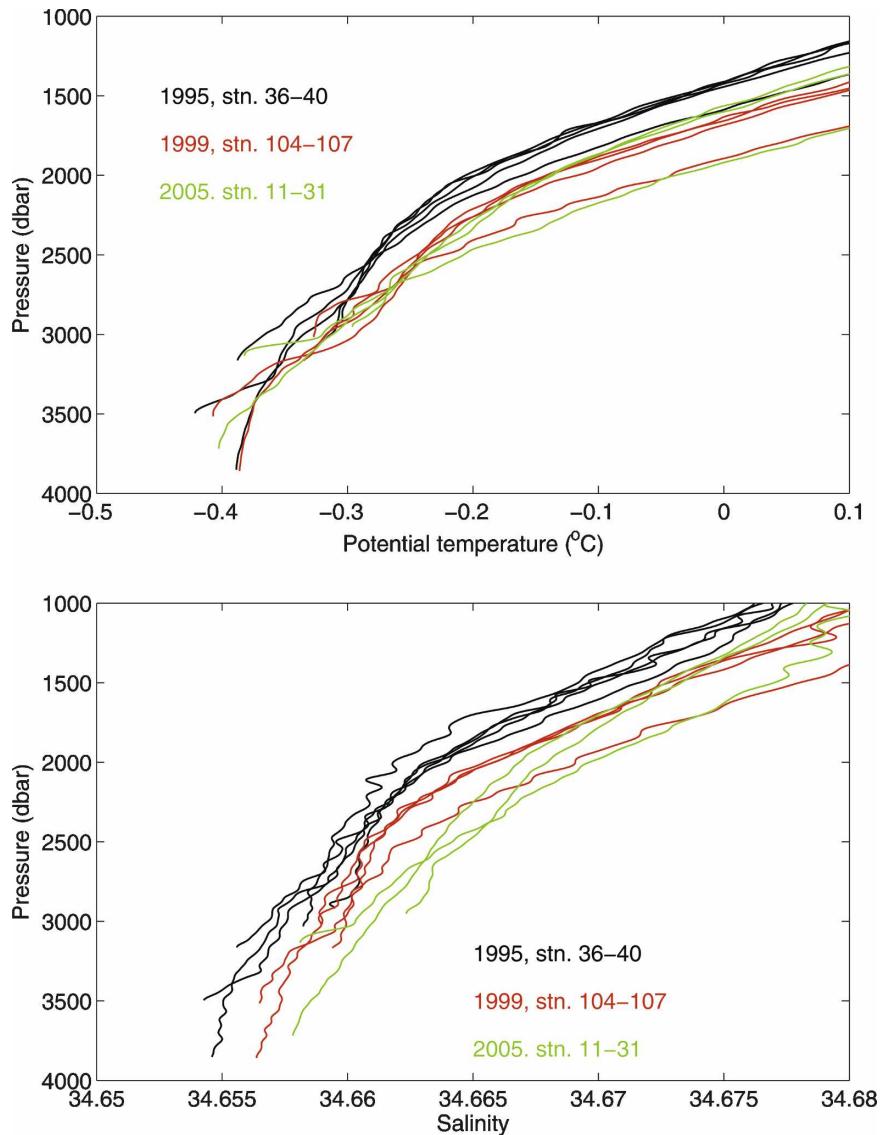


FIG. 8. (top) Potential temperature ($^{\circ}\text{C}$) and (bottom) salinity vs pressure for stations south of the Southern Boundary in 1995 (black), 1999 (red), and 2005 (green). Profiles have been smoothed with a 20-dbar Butterworth filter. Station positions are shown in Fig. 2.

South Sandwich Island arc (Meredith et al. 2001; Fig. 1). To reach these gaps, this water mass flows northward on the eastern side of the South Sandwich Islands, and its southward and westward entry to the Scotia Sea in 1999 contrasts with the 1995 and 2005 results, which indicate that the deepest layer of Lower WSDW had crossed the South Scotia Ridge into the Scotia Sea instead. The θ - S inflection is not evident north of the Southern Boundary in 1999, or south of the Southern Boundary in either of the other years. The 1999 pattern (an inflection south of the Southern Boundary but absent to the north) indicates that the Southern Boundary constitutes a dynamical barrier that plays an important

role in controlling the spreading of Lower WSDW into and within the Scotia Sea.

5. Discussion

a. Warm deep water

The pattern of θ and S variations in WDW in the Scotia Sea is qualitatively consistent with changes observed farther south in the Weddell Sea (Fahrbach et al. 2004; Robertson et al. 2002), with warming and salinification up to the mid- and late 1990s, followed by a subsequent cooling and freshening. These combined results are strongly suggestive of large-scale changes in

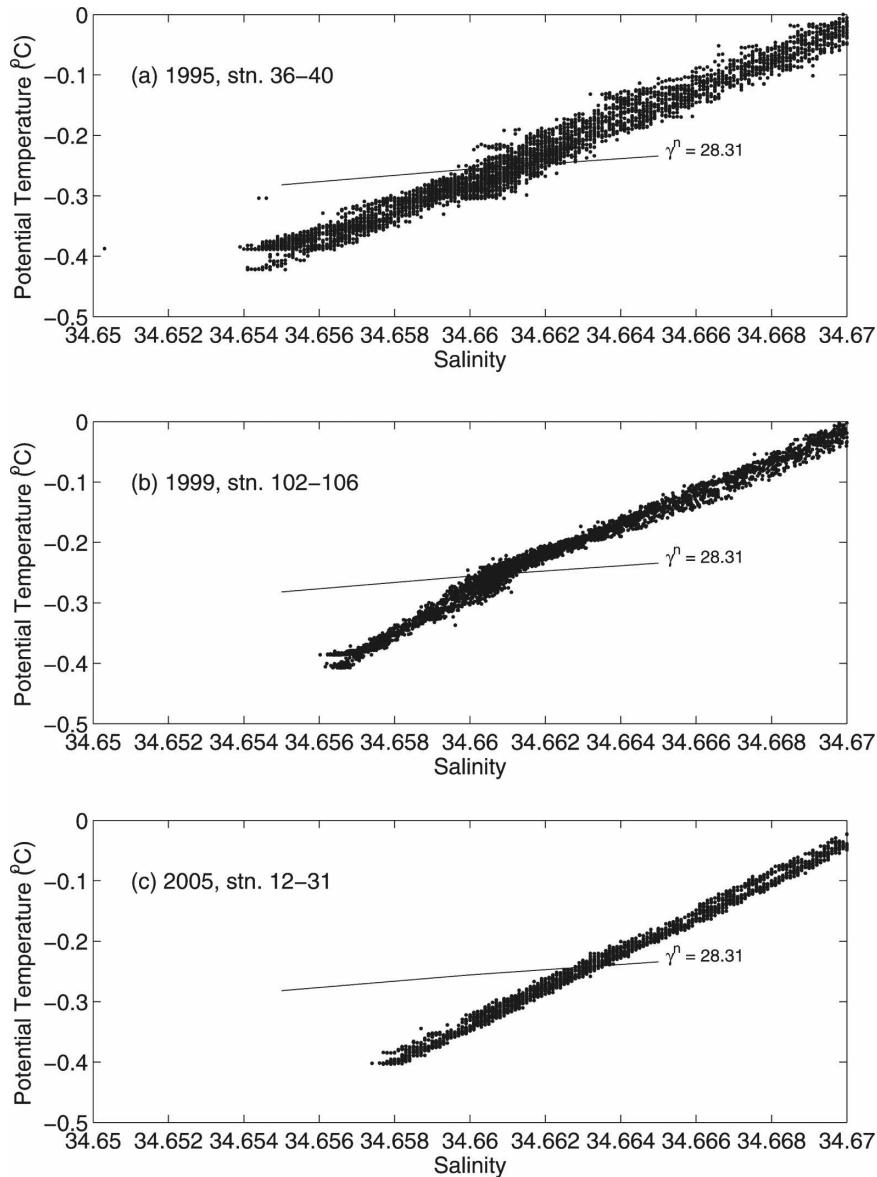


FIG. 9. Potential temperature–salinity curves for (top) 1995, (middle) 1999, and (bottom) 2005 for stations south of the Southern Boundary. An inflection at $\gamma^n = 28.31 \text{ kg m}^{-3}$, the Lower Weddell Sea Deep Water upper boundary, visible in 1999, is absent in 1995 and 2005.

the WDW properties in the Weddell Gyre being fed through to the Scotia Sea from upstream following the general circulation, with transit times of roughly a few years between the southern Weddell and Scotia Seas. We note that the $\sim 0.03^\circ\text{C}$ and ~ 0.002 increases in WDW θ and S values observed during the late 1990s by Fahrbach et al. (2004, their Fig. 5) at the Greenwich meridian in the Weddell Sea are distinctly less than the increases presented here between 1995 and 1999 in the Scotia Sea; similarly, the decrease in θ of $\sim 0.05^\circ\text{C}$ in the Weddell Sea after the turn of the century is less than

the cooling of WDW in the Scotia Sea between 1999 and 2005. However, the Weddell Sea figures are based on basin-wide averages that integrate some complex patterns of warming and cooling (Fahrbach et al. 2004); thus, the derived magnitudes of the changes here are expected to be smaller.

An analysis of WDW θ changes in “inflowing” and “outflowing” regions of the Weddell Gyre showed WDW θ peaking in 1995 in the inflowing region, with more recent values falling by around 0.2°C through to 1999 (Robertson et al. 2002). In the outflowing region,

TABLE 2. Area, volume, inflowing volume transport, and mean residence time of three overlapping density classes of WSDW with varying upper boundaries. Areas and volumes estimated by objective mapping of density profiles obtained from historical atlas data, the three sections used in this study, and six occupations of a transect across Drake Passage. Method described more extensively in Heywood et al. (2002). Uncertainties in the areas and volumes estimated ad hoc by varying objective mapping parameters. Volume transports over the South Scotia Ridge estimated from the box inverse model solution of Naveira Garabato et al. (2003). Errors correspond to one std dev uncertainties from the inverse model. Mean residence times calculated by dividing volumes by volume transports.

γ^n of upper water mass boundary (kg m^{-3})	Area in Scotia Sea (m^2)	Volume in Scotia Sea (m^3)	Inflowing volume transport (Sv; $1 \text{ Sv} \equiv 10^6 \text{ m}^3 \text{ s}^{-1}$)	Mean residence time (yr)
28.26	$(1.1 \pm 0.2) \times 10^{12}$	$(8.7 \pm 1.9) \times 10^{14}$	5.9 ± 1.7	4.7 ± 1.7
28.31	$(7.0 \pm 1.5) \times 10^{11}$	$(3.1 \pm 0.8) \times 10^{14}$	3.5 ± 1.2	2.8 ± 1.2
28.35	$(1.1 \pm 0.4) \times 10^{11}$	$(5.2 \pm 2.4) \times 10^{13}$	1.7 ± 0.5	1.0 ± 0.5

a large scatter in WDW θ values during the late 1990s made trends in this area hard to identify. The magnitudes of the changes in the WDW values in the Scotia Sea we present here are somewhat less than the long-term (~ 30 yr) warming trends seen in the WDW in the Weddell Sea (Robertson et al. 2002). This latter time series includes several instances of cooling within the general upward trend, and it is unclear whether the recent cooling seen in the Weddell Sea (Fahrback et al. 2004) is a reversal of the long-term warming trend or another instance of interannual variability against a background of continuing rising temperatures.

With regard to the timing of the peak WDW temperatures in the Weddell and Scotia Seas, the sparse temporal sampling in the latter precludes clear identification of the year of highest temperatures, and thus we cannot be precise concerning a lag between the two basins. Notwithstanding, the connection between WDW properties in the Weddell Sea and Scotia Sea established here implies that continued changes within the Weddell Sea are likely to be reflected as changes in the Scotia Sea in years to come.

b. Weddell Sea Deep Water

The changes identified in WSDW in the Scotia Sea are complex, and are summarized here by the following three main observations:

- 1) North of the Southern Boundary, WSDW warmed by around 0.05°C between 1995 and 1999, and again by the same amount between 1999 and 2005. A concurrent small salinification (0.005) occurred during the 1995–2005 period. The overall volume of WSDW in this part of the Scotia Sea decreased through the sequence of measurements.
- 2) South of the Southern Boundary, a comparable level of warming was seen between 1995 and 1999, but there was no significant warming between 1999

and 2005, and no significant salinity change was detected.

- 3) Within the Lower WSDW layer, changes in circulation were observed, with the densest waters crossing the South Scotia Ridge prior to the 1995 and 2005 sections, but entering the Scotia Sea from its eastern and northeastern flanks prior to 1999.

The observations presented here are best explained by a superposition of a large-scale decadal trend toward higher θ and S in the WSDW of the northern Weddell Sea, and shorter-period (interannual) variability in the export of the densest classes of WSDW to the Scotia Sea (see below). There are two factors that dictate the degrees to which these processes combine to control the observed properties of WSDW in the Scotia Sea, namely, the residence times of different classes of WSDW in the basin, and the Southern Boundary's role as a dynamical barrier to the northward spreading of this water mass.

The mean residence time of a particular class of WSDW in the Scotia Sea is the characteristic period over which the properties of inflowing waters are averaged. This period determines the time scales of variability that the class of WSDW exhibits, and hence indicates the degree to which the variability could be aliased by temporally sparse sampling. Mean residence times (Table 2) are calculated here by dividing the volume of each WSDW density class within the basin by the volume transport of that density class over the South Scotia Ridge (Heywood et al. 2002; Naveira Garabato et al. 2003). These calculations suggest that the overall WSDW layer in the Scotia Sea is renewed on average every ~ 5 yr, whereas the Lower WSDW sub-layer is flushed more rapidly (around every 3 yr). The densest classes of WSDW present in the Scotia Sea (here taken as $\gamma^n > 28.35 \text{ kg m}^{-3}$) are renewed on approximately annual time scales. These are found only south of the Southern Boundary, and the difference in

mean residence times for these classes compared with WSDW overall reflects the role of the Southern Boundary as a dynamical barrier to the northward flow of the water mass. The time intervals between the sections used here is comparable to the mean WSDW residence time in the Scotia Sea, but is significantly longer than that of the densest WSDW classes. The variability in denser WSDW properties is thus likely to be aliased within the hydrographic data, whereas the changes in WSDW characteristics north of the Southern Boundary will be well represented.

1) OBSERVATION: CONTINUOUS WARMING AND SALINIFICATION OF WSDW NORTH OF THE SOUTHERN BOUNDARY

Knowing the potential extent of aliasing associated with our sampling interval allows us to elucidate which of the two aforementioned processes (a large-scale decadal trend toward higher θ and S in the WSDW of the northern Weddell Sea, and interannual variability in the export of the densest classes of WSDW to the Scotia Sea) may explain each of our observed WSDW changes. The decadal trend alone may account for the first observation, because the residence time calculation indicates that the WSDW north of the Southern Boundary reflects long-term changes in the properties of the WSDW exported from the Weddell Sea.

Several previous studies (Andrié et al. 2003; Coles et al. 1996; Johnson and Doney 2006; Zenk and Morozov 2007) have documented a warming at a comparable rate in the WSDW in the South Atlantic as far north as the equator, consistent with the bulk warming of WSDW observed here “upstream” in the Scotia Sea. Decadal warming of the recently ventilated WSDW in the Weddell Sea has not been unambiguously demonstrated, because of the lack of a long time series of observations in the boundary currents directly upstream of the overflow regions at the South Scotia Ridge, but we note here that such a warming would be compatible with the observed decadal-scale warming of WSBW during the 1989–2003 period (Fahrbach et al. 2004). WSBW contains a significant proportion of recently ventilated shelf waters and is not directly exported from the Weddell Sea after formation, so changes within this layer are very possibly echoed by changes within the boundary current of WSDW (which contains a higher proportion of recently ventilated waters than the recirculated WSDW in the central Weddell Sea) that feeds the Scotia Sea.

As with the warming trend, a potential salinification of the recently ventilated WSDW in the northern Weddell Sea has not been shown unambiguously because of

the lack of appropriate observations. Robertson et al. (2002) were unable to deduce significant temperature trends for WSDW in the northwestern Weddell Sea because of large scatter in their data, and the same would appear to be true for salinity [e.g., the change of salinity we observe in the Scotia Sea, around 0.005, is small compared to the scatter shown in Fig. 9 of Robertson et al. (2002)].

2) OBSERVATION: CESSATION OF WARMING OF WSDW SOUTH OF THE SOUTHERN BOUNDARY AFTER 1999

Mean residence times for WSDW density classes that lie exclusively south of the Southern Boundary are of the order of 1 yr. Therefore, in the southern Scotia Sea, the impact of the decadal-scale warming trend discussed in 5b(1) is likely to be superposed on interannual variability in the WSDW overflows of the South Scotia Ridge. The mechanisms controlling the variability in these overflows were outlined by Meredith et al. (2001), and are illustrated here schematically (Fig. 10). It has been argued that changes in the intensity of the baroclinic component of the Weddell Gyre circulation (most likely associated with changes in integrated wind stress curl over the Weddell Sea) will affect the depth levels at which isopycnals intersect with the South Scotia Ridge, and hence determine the densest class of WSDW able to overflow into the Scotia Sea. A decrease (increase) in the density of the deepest layer of WSDW overflowing the ridge will result in a warming (cooling) on pressure surfaces within the Scotia Sea, but there is no discernable change on density surfaces. Accompanying S changes will occur, but of a magnitude comparable to the level of detectability. In the context of this conceptual model, a spinup of the Weddell Gyre between 1995 and 1999, followed by a spindown prior to 2005, would explain our pattern of observations, with the absence of significant changes south of the Southern Boundary between 1999 and 2005 resulting from the opposing effects of interannual variability associated with Weddell Gyre intensity (cooling) and longer-term (decadal) warming.

The extent to which the baroclinicity of the Weddell Gyre varies is poorly known, but there is some evidence that changes of the magnitude required to explain our observations occur. One piece is provided by a 6-yr time series of water mass properties collected with moored instrumentation immediately upstream of the Orkney Passage (Fig. 1) as part of the Consortium on the Ocean's Role in Climate (CORC) Abrupt Climate Change Studies (ARCHES) program (Visbeck et al. 2001). Although this mooring does not cover the den-

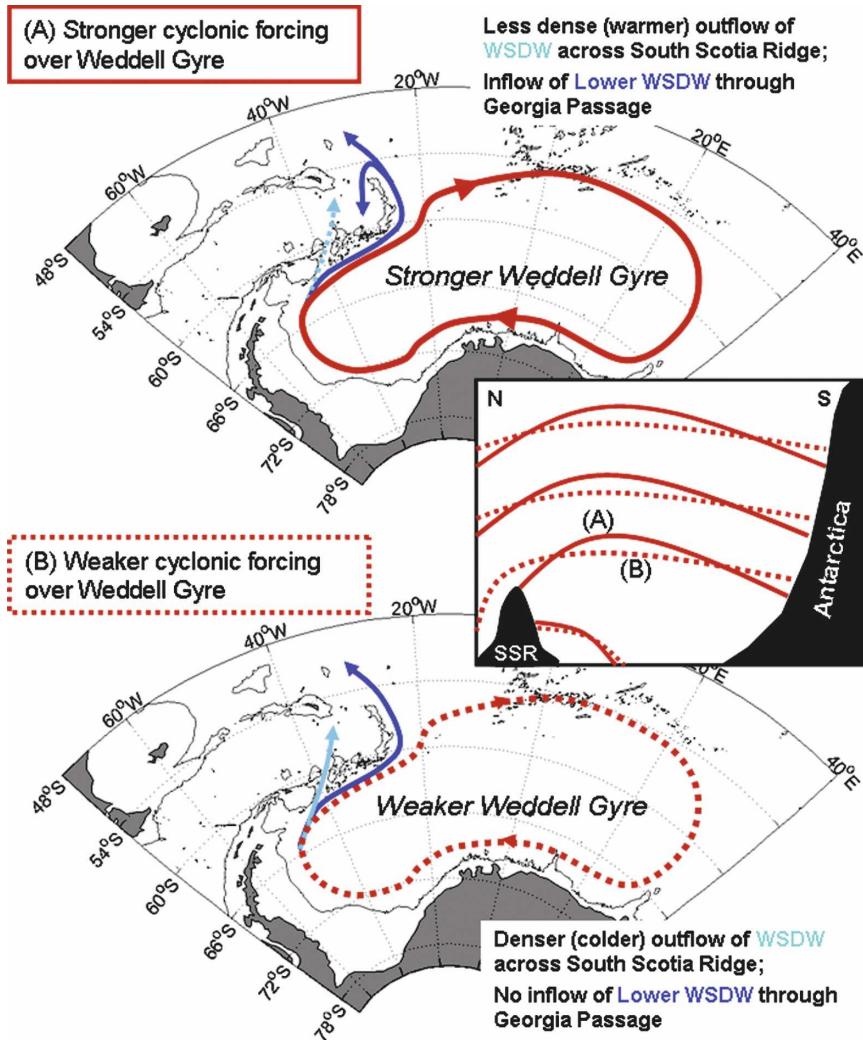


FIG. 10. Schematic of the impact of changing atmospheric cyclonic forcing on Weddell Gyre intensity and the consequences for the export of WSDW across the South Scotia Ridge (SSR) and the abyssal circulation of Lower WSDW. The inset (adapted from Coles et al. 1996) shows schematically the steepening (A) or slumping (B) of isopycnal surfaces in the Weddell Sea in response to gyre spinup/down, and the consequent change in the density of the deepest WSDW exported across the SSR to the Scotia Sea. (top) Case (A) is illustrated, where a stronger Weddell Gyre leads to less dense (warmer) WSDW crossing the SSR, and Lower WSDW entering the Scotia Sea through Georgia Passage and environs at the northeastern edge of the basin. (bottom) Case (B) is shown, where a weaker Weddell Gyre leads to denser (colder) WSDW crossing the SSR, and the absence of an inflow of Lower WSDW through Georgia Passage and environs.

sities that typify the WSDW in the Scotia Sea, and it is of shorter duration than the time period covered by our sections, it does provide useful information on the characteristic variability in density on a depth surface exhibited by WSDW in the boundary current of the northern Weddell Sea. The records show that at three different levels within 500 dbar of the seabed in water depths of around 3059 m, pronounced excursions of the order of 0.005 kg m^{-3} on interannual time scales

(around 1–3 yr) are prominent within the core of recently ventilated WSDW in the northern Weddell Gyre (Fig. 11). The interannual changes in density show a large degree of vertical coherence, consistent with these signals, resulting from variations in the baroclinicity of the gyre. Hydrographic profiles from this region indicate that a change in density at the mooring of the order of 0.005 kg m^{-3} corresponds to a change in isopycnal depth of around 100 m, with associated changes of ap-

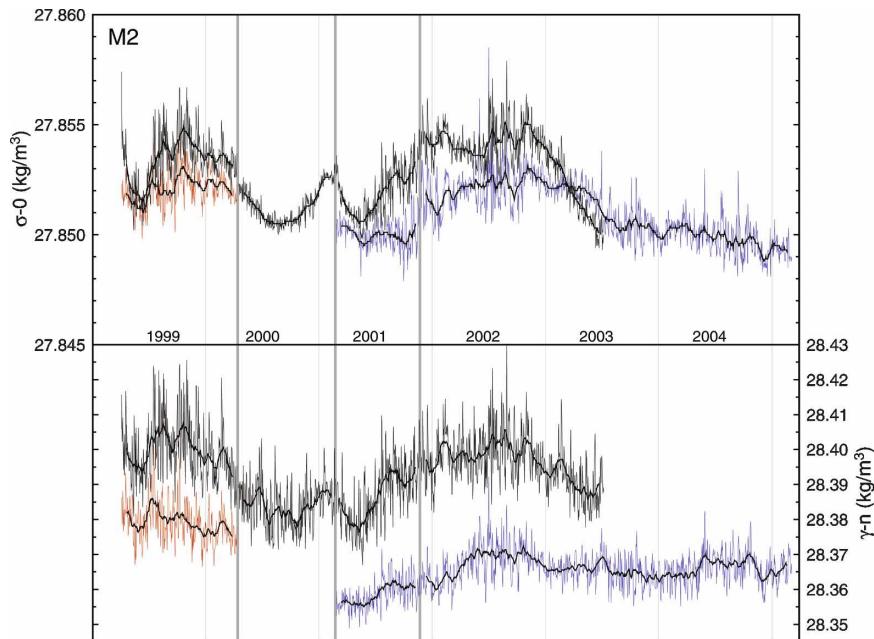


FIG. 11. Time series of (top) potential density (σ_0) and (bottom) neutral density (γ_n) at the CORC ARCHES M2 mooring site, color coded to indicate pressure separations of instruments from the seabed (red = 345 dbar; black 1999–2001 = 15 dbar; black 2001–03 = 31 dbar; blue = 478 dbar). Moorings deployed in 3059-m water depth (see Fig. 1). Pronounced excursions in density ($\sim 0.005 \text{ kg m}^{-3}$) on time scales of around 1–3 yr exhibit significant vertical coherence.

proximately 0.04°C in θ and 0.002 in S on a pressure surface. These changes are comparable with the variability in θ and S observed in the mooring data on interannual time scales (A. L. Gordon et al. 2008, unpublished manuscript). An interannual perturbation to the baroclinicity of the Weddell Gyre of the magnitude implied by the mooring density data would be sufficient to account for the cessation of the warming of the WSDW south of the Southern Boundary after 1999, while resulting in no detectable change in S , in agreement with our observations.

Further evidence for variability in the intensity of the baroclinic Weddell Gyre circulation is given by Martinson and Iannuzzi (2003), who use available hydrographic data from across the gyre to investigate periods of spinup and spindown during the last quarter of the twentieth century. These authors note significant changes in the gyre's baroclinicity on interannual time scales, which they relate to variations in the phase of the El Niño–Southern Oscillation (ENSO) phenomenon. During an El Niño event, the Weddell Gyre was deduced to have spun up in response to changing winds, and the converse was inferred in response to La Niña. If this relationship applies consistently, as Martinson and Iannuzzi (2003) propose, it might suggest that the apparently anomalous conditions that we observed in

the Scotia Sea during 1999 could be connected to the very strong El Niño event that occurred in 1997/98, allowing for a 1–2-yr lag for baroclinic spinup and for the anomalous WSDW that crosses the South Scotia Ridge to spread to the location of our section.

3) OBSERVATION: CHANGES IN LOWER WSDW CIRCULATION

The change in circulation of Lower WSDW between 1995 and 1999 (Meredith et al. 2001) was seen to subsequently reverse up to 2005. This pattern of change is consistent with observation 2 [section 5b(2)], and is similarly controlled by variations in the intensity of the Weddell Gyre (see Fig. 10). The increase in the density of the deepest WSDW overflowing the South Scotia Ridge associated with the baroclinic spindown of the Weddell Gyre after 1999 would have precluded the import of the more saline variety of Lower WSDW through Georgia Passage and gaps in the South Sandwich Arc. (A denser outflow through these gaps would act to block the southwestward flow of Lower WSDW through Georgia Passage into the Scotia Sea abyss). Consequently, 1999 was the only year of the three that showed the influence of the more saline type of Lower WSDW, resulting from the reversals in the direction of

abyssal flow at the northeastern and eastern edges of the Scotia Sea before and after this date.

6. Conclusions

A hydrographic section in the eastern Scotia Sea of the Southern Ocean was occupied three times (in April 1995, April 1999, and February 2005), and significant changes in water mass properties were observed. The WDW south of the Southern Boundary warmed and salinified between 1995 and 1999, and then cooled and freshened through to 2005. These changes appear to be the consequence of decadal-scale variability in the properties of this water mass upstream in the Weddell Sea, propagated into the Scotia Sea via the Weddell Gyre. The likely influence of the Weddell Polynya on WDW properties has been highlighted (Robertson et al. 2002), and it has been argued that this event may have been triggered by anomalous atmospheric forcing associated with the Southern Hemisphere Annular Mode, and possibly linked to a La Niña event (Gordon et al. 2007). In the context of these studies, our results indicate that climate-scale forcing over the Weddell Sea has impacts that are seen outside the region because the induced variations are propagated northward.

We have reported several changes to the WSDW flowing through the Scotia Sea. North of the Southern Boundary, WSDW warmed and salinified between 1995 and 2005, whereas south of the Southern Boundary the warming was only observed through to 1999. Reversals in the abyssal circulation of the densest classes of WSDW were also observed. The observed WSDW changes are best explained by a superposition of a large-scale decadal trend toward higher θ and S in the WSDW of the northern Weddell Sea, and shorter-period (interannual) variability in the export of the densest classes of WSDW to the Scotia Sea.

We have argued that the interannual variability in the export of WSDW to the Scotia Sea is likely related to changes in the baroclinicity of the Weddell Gyre. These changes may be controlled by cyclonic atmospheric forcing over the gyre, which has been shown to be modulated by ENSO on interannual time scales (see Martinson and Iannuzzi 2003, and references therein). We conjecture that if the interannual modulation of WSDW export by ENSO was also active on longer (decadal) time scales (as implied by the work of Martinson and Iannuzzi 2003), it may contribute to the WSDW warming observed throughout the South Atlantic. ENSO has undergone long-period changes in recent decades, with a higher preponderance of El Niño events since the 1970s; this would enhance the Weddell

Gyre's baroclinicity, and hinder the export of the coldest WSDW.

The WDW and WSDW changes that we observe have potential climatic consequences that transcend regional scales. For example, if the decrease in volume of WSDW in the Scotia Sea between 1995 and 2005 is indicative of a longer-period trend, this will have implications for circulation and climate in much of the World Ocean. Warming of the abyssal layer has been reported for recent decades as far north as the equator in the western Atlantic Ocean. This warming is evidence of a major change in the properties of the lower limb of the oceanic overturning circulation, originating in the Weddell Sea and transiting through the Scotia Sea.

To put some of these changes in perspective, climate models suggest that the earth's energy imbalance is currently about $0.85 (\pm 0.15) \text{ W m}^{-2}$, with the oceans absorbing a large fraction of that heat gain (Hansen et al. 2005). The evolution of the heat content of WDW in the eastern Scotia Sea is equivalent to a local heat gain of $\sim 3 \text{ W m}^{-2}$ between 1995 and 1999, and a similar local heat loss through to 2005 [cf. the local heat gain of 4 W m^{-2} implied by the warming of the Weddell Sea between 1977 and 2001; Smedsrud (2005)]. In turn, the temperature increase of the WSDW layer in the eastern Scotia Sea between 1995 and 2005 associated with the decadal warming trend yields a local heat gain of $\sim 1 \text{ W m}^{-2}$. If the vertical profile of the warming rate below 3000 m between 60°S and the equator in the western basins of the South Atlantic for recent decades (Johnson and Doney 2006) was typical of the entire global ocean abyss, it would account for $\sim 0.2 \text{ W m}^{-2}$ of the earth's energy imbalance. Abyssal Pacific warming rates are smaller (Johnson et al. 2007), nonetheless abyssal water temperature variations such as those presented here are not only of high regional importance, but also may contribute significantly to the global heat budget.

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